

EVALUATION OF THE SEISMIC RESPONSE USING THE RANDOM VIBRATION THEORY IN THREE SOIL PROFILES OF LIMA, PERU

EVALUACIÓN DE LA RESPUESTA SÍSMICA UTILIZANDO LA TEORÍA DE VIBRACIONES ALEATORIAS EN TRES PERFILES DE SUELO DE LIMA, PERÚ

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ABSTRACT

In general, seismic response analyses are performed using a minimum number of seismic records as input motions in order to achieve a statistically strong estimation. Unfortunately, the available information recorded from the current seismic networks is still scarce regarding events with considerable magnitude. In this context, the Random Vibration Theory (RVT) arises as an alternative tool for performing site response analyses without the need of seismic records, since it only requires adequate probabilistic seismic hazard assessment.

In this study, RVT was applied to three shear-wave velocity profiles in Lima city with distinct geomorphological origin. These profiles are characteristic for gravelly, sandy and fine deposits so the influence of each soil type in their corresponding transfer function was taken into account. In that sense, the three RVT-based normalized response spectra show good agreement with the design spectra specified in the Peruvian code, despite some amplification in the short (below 0.10 s) and long (above 0.80 s) period ranges related to noise or far-field effects. Furthermore, RVT-based response spectra for La Punta and Villa El Salvador districts show good agreement with the time-series based analyses from a previous study.

In addition, spectral acceleration values surpass those specified in the Peruvian code for a range beyond the corner period. This could suggest that the soil profile characterization based on the time-averaged shear wave velocity from the upper 30 m might be insufficient to evaluate the overall seismic behavior of a soil deposit. Therefore, additional parameters that account for the deeper soil substructure might be required.

Keywords: Site Response Analysis; Probabilistic seismic hazard analysis; Shear-wave velocity profiles; Random Vibration Theory; Lima

RESUMEN

En general, los análisis de respuesta sísmica se realizan utilizando un número mínimo de registros sísmicos como movimientos de entrada para lograr una estimación estadísticamente sólida. Desafortunadamente, la información disponible registrada a partir de las redes sísmicas actuales es aún escasa respecto a eventos de magnitud considerable. En este contexto, la Teoría de Vibraciones Aleatorias (RVT) surge como una herramienta alternativa para realizar análisis de respuesta de sitio sin necesidad de registros sísmicos, ya que solo requiere de una adecuada evaluación probabilística de la amenaza sísmica.

En este estudio, la RVT se aplicó a tres perfiles de velocidad de ondas de corte en la ciudad de Lima con distinto origen geomorfológico. Estos perfiles son característicos para depósitos gravosos, arenosos y finos por lo que se tuvo en cuenta la influencia de cada tipo de suelo en su correspondiente función de transferencia. En ese sentido, los tres espectros de respuesta normalizados basados en RVT muestran una buena concordancia con los espectros de diseño especificados en el código peruano, a pesar de cierta amplificación en los rangos de período corto (inferior a 0,10 s) y largo (superior a 0,80 s) relacionados con el ruido o con el tiempo. -efectos de campo. Además, los espectros de respuesta basados en RVT para los distritos de La Punta y Villa El Salvador muestran una buena concordancia con los análisis basados en series temporales de un estudio anterior.

Además, los valores de aceleración espectral superan los especificados en el código peruano para un rango más allá del período de esquina. Esto podría sugerir que la caracterización del perfil del suelo basada en la velocidad de la onda de corte promediada en el tiempo desde los 30 m superiores podría ser insuficiente para evaluar el comportamiento sísmico general de un depósito de suelo. Por lo tanto, podrían ser necesarios parámetros adicionales que tengan en cuenta la subestructura del suelo más profundo.

Palabras clave: Análisis de respuesta del sitio, Análisis probabilístico de peligrosidad sísmica, Perfiles de velocidad de onda cortante, teoría de la vibración aleatoria, lima

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1. INTRODUCTION

Seismic response analysis is one of the most common problems in earthquake geotechnical engineering and it is especially important for sites with high seismic activity, such as countries located in the so-called Pacific Fire Ring, as is the case of Peru. In general, site response analyses are performed considering a propagation of a predefined input motion from the engineering bedrock through the overlying soil, in order to obtain a surface intensity index. This procedure allows the prediction of ground surface motions characteristics for a predefined level of seismic hazard expected in the site of study.

Typically, the required input motion is obtained from processing a set of selected accelerograms to match the Uniform Hazard Spectrum (UHS) defined from a seismic hazard analysis performed for the site under study. This method, hereinafter referred to as time-series, requires to have a minimum number of records to achieve a statistically strong estimation. Unfortunately, CISMID seismic network (REDACIS), as well as others deployed in Lima city, have so far generated scarce seismic information related to events with considerable values of moment magnitude.

In this context, the Random Vibration Theory (RVT) arises as an alternative method for seismic response analyses that can be performed by only requiring the UHS from Probabilistic Seismic Hazard Analysis (PSHA) and a strong motion duration.

RVT was originally applied in seismology by Hanks and McGuire[1] to predict peak ground motion parameters based on the earthquake magnitude and site-to-source distance, and further developed by Boore[2] and Boore & Joyner[3]. Schneider et al.[4] applied RVT to site response analysis to simulate site-specific acceleration response spectra by defining the input motion based on a seismological source theory to be later convolved. Thus, Fourier amplitude (FAS) and response spectra could be estimated by means of RVT. Later studies, e.g. Silva et al.[5] and Rathje & Ozbey[6], validated RVT procedure for seismic response analysis, although it requires seismological parameters to be carefully chosen in order to be consistent with the site-specific seismic hazard at rock level.

To overcome the need of adjusting multiple seismological parameters, Rathje & Ozbey[6] developed the Inverse Random Vibration Theory (IRVT) to convert an acceleration response spectrum into a frequency domain FAS, which is a simpler form to define the input motion for the seismic analysis using RVT. The application of RVT has dramatically increased in time. For example, the nuclear industry has been switching to the RVT-based approach to

perform the seismic response analysis as an alternative to the time-series method over the last years.

The first approach to RVT application in earthquake engineering in Peru was carried out by Chavez[7]. In which the response spectra from a batch of recorded seismic events were obtained. In addition, theoretical response spectra, based on the RVT and the omega square (ω^2) seismological model[8][9], were computed. Chavez[7] also performed a sensitivity analysis of the seismological parameters from the single corner frequency model to maximize the fit between the recorded and the theoretical response minimizing the standard deviation obtained through the Bayesian linear regression.

In this context, the objective of this research is to perform site response analysis using the RVT methodology, incorporated in the software Strata[10], on three representative and well differentiated soil profiles in Lima. Reliability of RVT is performed by the comparison of the spectral shape of the results with the pseudo-acceleration spectra from signals recorded in stations nearby the studied sites, for moderate to strong seismic events, and those from previous site response studies. Furthermore, sensitivity analyses, with respect to the duration of the seismic record and the shear-wave velocity (V_s) of the engineering bedrock, are also evaluated.

2. METHODOLOGY

2.1. Probabilistic seismic hazard analysis

Uniform Hazard Spectrum (UHS) and seismic disaggregation were obtained from the study of seismogenic regions, recurrence laws, attenuation laws and their ponderation proposed by Roncal[11]. Seismic response analysis included a sensitivity evaluation to consider the influence when varying the stiffness of the engineering bedrock for two cases: soil type A ($V_s > 1500$ m/s) and type B ($V_s > 760$ m/s), according to ASCE 7-16[12]. Consequently, PSHA and seismic disaggregation were performed for each of these cases for a 10% probability of exceedance of a seismic event of 500 years return period. In Figure 1, UHS obtained from the PSHA are presented. As expected, the acceleration spectrum was slightly higher for the type B soil in a narrow period range so it can be considered that there was no significant difference in the seismic scenario identified as critical from the seismic disaggregation.

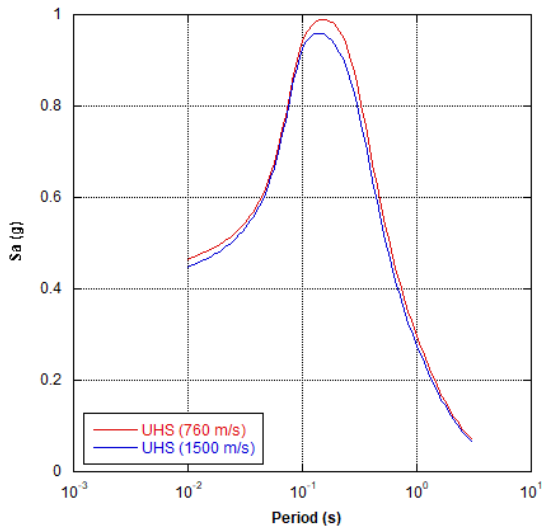


Fig. 1 Uniform Hazard Spectrum for soil type A and soil type B.

The urban area of Lima city is mainly composed of one to five story buildings that are typically characterized by fundamental periods that range from 0.05 s to 0.50 s. For this range of interest, the seismic scenario with the highest contribution to the seismic hazard is represented by an interface scenario of 50 km closest rupture distance (R_{rup}), 90 km hypocentral distance (R_{hypo}) and 7.9 seismic moment magnitude (M_w).

2.2. Ground motion duration

Ground motion duration is an important input for the RVT-based seismic response analysis. For this research a sensitivity analysis was performed considering two approaches. First, duration was defined based on the ground motion model given by Campbell[13], which includes source and path duration. On the other hand, a second procedure was considered, based on the attenuation law developed by Boroschek et al.[14].

2.2.1. Ground motion duration model

Ground motion duration includes two terms: source (T_0) and path (T_r) durations, as

$$T_d = T_0 + T_r \tag{1}$$

Source duration (T_0) is the inverse of the corner frequency (f_0). If the seismic source is assumed to come from a point source, the Brune model[8] can be used to calculate the corner frequency as follows

$$f_0 = 4.9 \times 10^6 \beta_0 \left(\frac{\Delta\sigma}{M_0} \right)^{\frac{1}{3}} \tag{2}$$

where β_0 is the shear wave velocity of the crust in km/s, $\Delta\sigma$ is the stress drop in bars and M_0 is the

seismic moment in dinas.cm. Chavez[7], based on a sensitivity analysis and comparison of recorded and calculated response spectra, concluded that the seismological parameters that best fit the Peruvian seismic environment are $\beta_0 = 3.9$ km/s and $\Delta\sigma = 185$ bars.

The seismic moment correlates with the seismic moment magnitude (M_w) under equation (3). As stated by Rathje et al. [19], the single corner frequency model has produced accurate results.

$$M_0 = 10^{1.5M_w + 16.05} \tag{3}$$

Path duration depends on the hypocentral distance (R_{hypo}) and can be estimated from

$$T_r = 0.05 \times R_{hypo} \tag{4}$$

where R_{hypo} is defined in km.

The duration for the seismic scenario that provides the highest contribution to the seismic hazard in the range of periods of interest ($M_w = 7.9$ and $R_{hypo} = 90$ km) is 15 s.

2.2.2. Ground motion duration attenuation law

Boroschek et al.[14] developed an attenuation law for strong ground motion duration based on a dataset of seismic records in Chile. This duration is defined as the time lapse in which energy, within the range of 5% and 95% based on the Arias intensity, is released. The formal function depends on the seismic scenario pair (closest rupture distance and moment magnitude), as presented as follows

$$\log D = c_1 + c_2 M_w + c_3 \ln \ln R + c_4 F_{soil} + c_5 F_{inter} + c_6 F_{intra} \tag{5}$$

where F_{soil} , F_{inter} and F_{intra} are binary variables that are defined based on the site conditions and seismic scenario. The coefficients are summarized in Table I.

For the study site, the evaluation was performed for the engineering bedrock and the seismic scenario pair (M_w and R_{hypo}) of the interface event that provides the highest contribution to seismic hazard level. Based on this approach, the duration for the strong motion resulted as 24 s.

TABLE I
Formal function coefficients for the strong motion duration attenuation law (Boroschek et al. 2019)

C1	C2	C3	C4	C5	C6
-1.052	0.173	0.454	0.031	1.081	0.908

2.3. Random vibration theory

RVT uses the Parseval theorem and Extreme Values Statistics (EVS) to predict the maximum acceleration in the time domain from the Fourier Amplitude Spectrum (FAS) and duration of the considered motion. The following is a summarized discussion about RVT. Further details can be found in [15] and [16].

Parseval's theorem states that the power of a seismic signal is conserved in frequency and time domain. Hence, total power can be computed in either domain. The root mean square acceleration (a_{rms}) is equal to the total power in the time history divided by the ground motion duration (T_d) and can be computed from the square of the FAS ($|A(f)|$) as presented in

$$a_{rms}^2 = \frac{1}{T_d} \int_0^{T_d} a^2(t) dt = \frac{2}{T_d} \int_0^\infty |A(f)|^2 df \quad (6)$$

RVT defines the peak factor (pf) as the ratio between a_{rms} and the peak acceleration, expressed as follows

$$a_{max} = pf \cdot a_{rms} \quad (7)$$

pf is calculated based on EVS as a function of the statistical moments of the FAS and the duration of the motion. Ocean wave amplitudes were studied by Cartwright and Longuet-Higgins[17] considering the probability distribution of maxima amplitudes to develop expressions for pf in terms of the signal characteristics. Later, Boore[18] derived an integral expression for the expected value of the peak factor in terms of the number of extrema (N_e), the number of zero crossing (N_z) and the bandwidth ($\xi = N_z/N_e$), which is expressed as

$$pf = \sqrt{2} \int_0^\infty \left\{ 1 - \left[1 - \xi e^{-z^2} \right]^{N_e} \right\} dz \quad (8)$$

2.4. Inverse random vibration theory (IRVT)

IRVT converts an acceleration spectrum to a Fourier Amplitude Spectrum (FAS) in the frequency domain. This process is not as straightforward as RVT. In order to achieve this inverse process, there are two limitations: the spectral accelerations are influenced by a range of frequencies and the peak factor cannot be calculated from the very beginning since it requires a Fourier Amplitude Spectrum.

In RVT, the peak spectral acceleration (S_a) is related to the rms spectral acceleration ($S_{a_{rms}}$) by the peak factor.

$$S_a^2 = pf_2^2 (S_{a_{rms}})^2 \quad (9)$$

To determine the $S_{a_{rms}}$ for a SDOF system, Parseval's theorem is used

$$S_{a_{rms}}^2 = \frac{2}{T_d} \int_0^\infty |A(f)|^2 |H_{f_n}(f)|^2 df \quad (10)$$

where $|H_{f_n}(f)|$ is the transfer function of a SDOF oscillator with natural frequency (f_n) and critical damping ratio ξ .

A first limitation is observed in (9). Spectral acceleration is affected by a range of frequencies in the FAS; thus, the estimation is not straightforward. However, some properties of the SDOF transfer function can be used to overcome this limitation. Typical transfer function of a SDOF has a narrow band for lightly damped systems and goes near to a value of 1 for low frequencies. In addition, large amplification exists near the natural frequency and tends to zero for higher frequencies. Using these properties, an approximation to the integral in (9) can be presented

$$S_{a_{rms}}^2 = \frac{2}{T_d} \left[\int_0^{f_n} |A(f)|^2 df + |A(f)|^2 \int_0^\infty |H_{f_n}(f)|^2 df - |A(f)|^2 f_n \right] \quad (11)$$

The transfer function integral is constant for a given natural frequency and damping ratio, allowing to simplify (10) as follows

$$|A(f_n)|^2 \approx \frac{1}{f_n \left(\frac{\pi}{4} - 1\right)} \left(\frac{T_d S_a^2}{2 pf^2} - \int_0^{f_n} |A(f)|^2 df \right) \quad (12)$$

The second limitation is solved by assuming the peak factor to estimate the initial Fourier Amplitude Spectrum. Typically, the iterative process is initiated by assuming a pf of 2.5. Further discussion regarding the methodology is included in [16] and [19].

RVT and IRVT procedures are included in the Strata software, developed by Rathje and Kottke[10], and were utilized in this study.

2.5. Seismic response analysis

One-dimensional wave propagation can be computed in the frequency domain using transfer functions for multi-layered soil deposits or in the time domain using multiple-degree-of-freedom (MDOF) dynamic response analysis. Due to the nature of the RVT-based analysis, convolution must be performed in the frequency domain, as no phase content is defined through this method. For the convolution, the equivalent linear analysis (EQL) is performed to account for the nonlinear behavior of the soil. This formulation is extensively discussed in [20].

Seismic response analysis is performed specifying an input motion at the base of the soil deposit defined as the calculated FAS of the acceleration spectrum for the engineering bedrock using the IRVT. Then, the input motion is propagated through the soil profile using the EQL formulation as implemented, for example, in Deepsoil[21] software. Finally, FAS at surface obtained from the EQL analysis is converted into acceleration spectrum using RVT. The aforementioned is presented in the workflow in Figure 2.

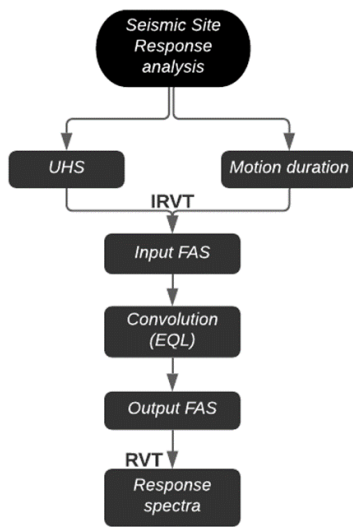


Fig.2 Workflow of the RVT-based seismic response analysis

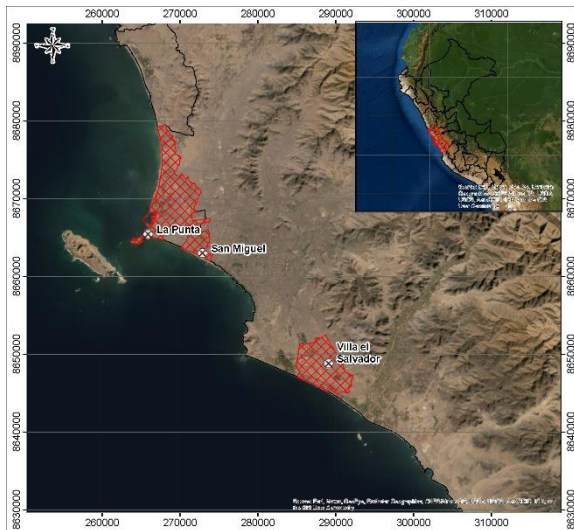


Fig. 3 Sites evaluated using the RVT-based seismic response analysis

TABLE II
Location of the evaluated sites

Site	Reference	Latitude (°)	Longitude (°)
San Miguel	“La Sagrada Familia” Park	-12.084	-77.086
Villa El Salvador	School 6066	-12.213	-76.939
La Punta	Direction of Hydrograph	-12.064	-77.155

3. Site characterization

Location of the evaluated sites is summarized in Table II and Figure 3.

The characterization of each profile was based on available information and complementary field surveys performed as part of this research.

3.1. Collected data

Collected data comprises previous studies, explorations (direct and indirect tests) and laboratory tests. In-situ explorations gathered data such as soil pits, boreholes and geophysical tests (active and passive source).

La Punta profile was adapted from the characterization of the DHN profile developed by Soto[22] which was based on two drill holes located near the study site[23] and the dynamic small strain properties from MASW tests[22] for the surficial layers and a microtremor array[24] for the deeper part.

Similarly, Villa El Salvador profile was also adapted from the characterization of the profile developed by Soto[22]. The profile was geotechnically interpreted mainly from a drillhole[25] and a microtremor array[24]. To complete the dispersion curve obtained in Calderon[24], complementary microtremor array measurements were conducted within the framework of this study which are further discussed in the field survey section.

San Miguel profile was characterized based on two boreholes located near the study site and a test pit[23], and the small-strain dynamic properties were inferred from a MASW[23] for surficial layers. The small-strain dynamic properties for the deeper part were interpreted from microtremor arrays carried out as part of this research and of Sencico project[26]. This survey is also further discussed in the following section.

3.2. Field surveys

Field surveys included microtremors array tests performed in San Miguel and Villa El Salvador. Two

types of sensors were used: the moving-coil velocimeter CR 4.5-1A and the servo velocimeter CV-374 AV2, manufactured by Anet Co., Ltd. and Tokyo Sokushin Co., Ltd., respectively. Microtremor arrays in San Miguel included linear and circular shapes which allowed us to interpret the dispersion curve of surface waves up to a period of 0.50 s. Larger microtremor arrays[26] were performed to complete the dispersion curve up to a period value of 0.70 s.

Circular microtremor arrays in Villa El Salvador adequately completed the dispersion curve obtained by Calderon et al.[24] within the range of 0.50 and 1.0 s. Good overlapping was observed between the measured and the reference dispersion curve assuring the reliability in the estimated shear-wave velocity values for deeper layers. Information regarding size and sampling frequencies for the conducted arrays are summarized in Table III. Figure 4 presents the sensor distribution for San Miguel and the location of a vertex sensor for Villa El Salvador.



Fig. 4 Field survey photographs. a) San Miguel 20 m array configuration. b) Villa El Salvador vertex sensor for the 200 m-array

TABLE III
Information about the conducted microtremor array measurements

Site	Sensor	Array type	Spacing/radius (m)	Sampling frequency (Hz)
San Miguel	CV-374 AV2	Circular	5	1000
			10	1000
			45	200
	CR 4.5-1A	Linear	0.5	2000
			2	2000
		Circular	5	1000
			10	1000
Villa El Salvador	CV-374 AV2	Circular	20	500
			45	500
			100	200
			200	200
			400	200

3.3. Sites profiles

Information about the estimated profiles in San Miguel, Villa El Salvador and La Punta is summarized in Table IV, Table V and Table VI, respectively. Figure 5 shows the shear wave velocity profiles and the horizontal-to-vertical spectral ratio from microtremors (mHVSR) for each site. It is important to highlight that the peaks in the long period range are evident for each profile, as these are representative of the dynamic behavior of the deep soil profiles that are potentially not captured by the soil classification system based on the Vs30 only.

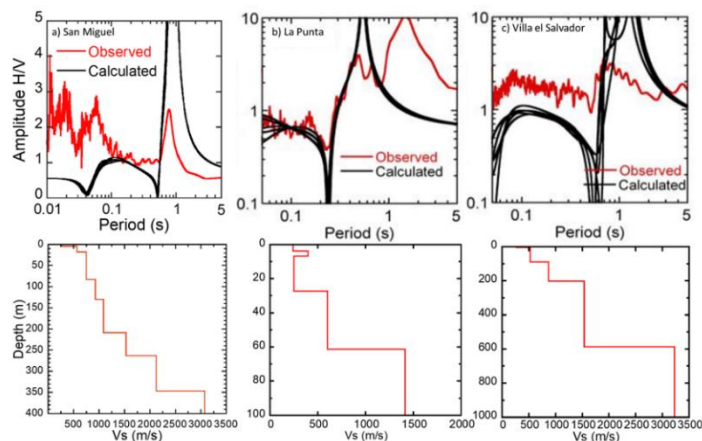


Fig. 5 Evaluated sites inverted profiles in San Miguel[26], La Punta[24] and Villa El Salvador[24]

TABLE IV
San Miguel geotechnical profile

Nro	Layer	Thickness (m)	Density (kN/m ³)	Average shear wave velocity (m/s)	Input motion applied in layer
1	Clay	2.3	18.0	289	
2	Gravel	24.5	22.0	585	
3	Gravel	56.2	22.0	746	
4	Gravel	47.4	22.0	929	x
5	Gravel	78.8	22.0	1089	
6	Gravel	-	22.0	1532	x

TABLE V
Villa El Salvador geotechnical profile

Nro	Layer	Thickness (m)	Density (kN/m ³)	Average shear wave velocity (m/s)	Input motion applied in layer
1	Sand	3.0	18.7	150	
2	Sand	3.0	18.7	320	
3	Sand	83.2	20.0	530	
4	Sand	112.5	20.0	870	x
5	Sand	-	20.0	1535	x

TABLE VI
La Punta geotechnical profile

Nro	Layer	Thickness (m)	Density (kN/m ³)	Average shear wave velocity (m/s)	Input motion applied in layer
1	Clay	1.1	18.7	189	
2	Gravel	19.8	22.0	336	
3	Clay	5.9	19.5	416	
4	Sand	9.6	19.5	431	
5	Gravel	6.0	20.5	602	
6	Clay	17.9	20.5	602	
7	Gravel	-	22	1414	x

Seismic analyses were carried out considering the dynamic properties degradation curves proposed by Menq[27] for gravelly and sandy soils, whereas Zhang model[28] was used for clayey soils. Based on

laboratory tests and technical literature available, index properties were selected for sandy ($C_u = 2.2$, $D_{50} = 0.17$ mm[25]), gravelly ($C_u = 10$ [22][29], $D_{50} = 30$ mm[22][29][30]) and clayey soils ($PI = 10$ [22][25]). Figure 6 presents the shear modulus degradation and damping curves defined for analyses.

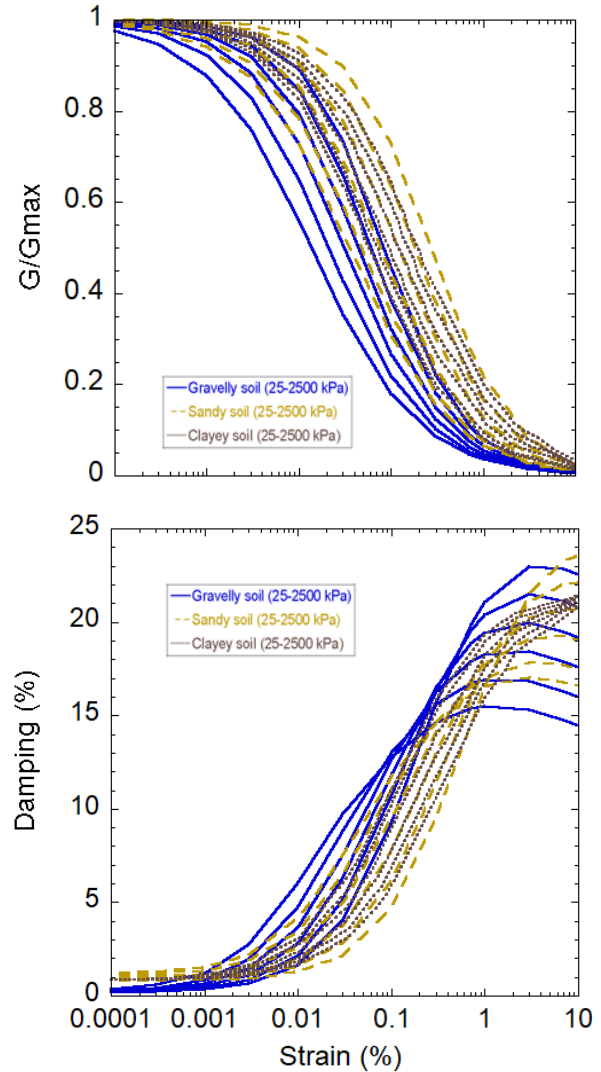


Fig. 6. Shear modulus degradation and damping curves evaluated in the seismic response analysis

4. RESULTS

For each site, acceleration spectra from the RVT-based seismic response analysis were compared with the design spectra from the Peruvian seismic design standard[31]. In the case of Villa El Salvador and La Punta, additional comparison was made considering the acceleration spectra from a previously developed study based on time series analysis[22]. It is noted that the UHS used in this research has larger amplitudes than the considered in Soto[22]; thus, the comparison is mainly focused on the spectral shape,

as the amplifications in this research are expected to be higher. Furthermore, seismic events with distinct

TABLE VII
Seismic records evaluated for the comparison and validation of RVT-based seismic response analysis

Site	Profile classification [31]	Station ID (Network)	Latitude (°)	Longitude (°)	Event Date	Epicenter	Magnitude	Epicentral distance (km)	Depth (km)
San Miguel	S1 ($V_{s30}=550$ m/s) ($500 < V_{s30} < 1500$ m/s)	PUCP (PUCP)	-77.08	-12.07	22/06/2021	Mala	MI 5.8	79	32
					25/08/2021	Callao	MI 5.1	55	32
					26/05/2019	Lagunas	Mw 8.0	723	135
Villa El Salvador	S2 ($V_{s30}=460$ m/s) ($180 < V_{s30} < 500$ m/s)	UNTELS (CIP)	-12.213	-76.932	22/06/2021	Mala	MI 5.8	61.7	32
		VSV (CISMID)	-12.213	-76.938	25/11/2013	Mala	MI 5.8	78.2	59
		VSAL (IGP)	-12.2	-76.9	26/05/2019	Lagunas	Mw 8.0	738	135
La Punta	S2 ($V_{s30}=390$ m/s) ($180 < V_{s30} < 500$ m/s)	DHN (CISMID)	-12.066	-77.156	22/06/2021	Mala	MI 5.8	82	32
					25/08/2021	Callao	MI 5.1	50	41
					26/05/2019	Lagunas	Mw 8.0	724	135
					15/08/2007	Pisco	Mw 7.9	182	40
		D855 (CIP)	-12.071	-77.164	22/06/2021	Mala	MI 5.8	82.5	32
		DSMI (CISMID)	-12.0608	-77.1449	22/06/2021	Mala	MI 5.8	82	32

epicentral distances and recorded in stations located nearby the evaluated sites were also used for comparing shapes of the RVT-based normalized spectra. The decision was made since the analyzed earthquakes produced responses significantly lower than an event for 500 years return period with 10% probability of exceedance.

Table VII summarizes the profile classification of each site according to the Peruvian standard [31], details regarding the stations near the evaluated sites and the recorded seismic events.

San Miguel RVT-based spectrum performed for engineering bedrock type B fits adequately under the design spectrum. Conversely, the spectrum for type A shows a slight difference around 0.80 s (Figure 7.a). This difference might be due to the influence of the deeper soil substructure in the amplification around the fundamental period, which is not taken into account when considering type B bedrock. However, this difference correlates well with the behavior found in seismic records for intermediate- and far-field earthquakes, such as the MI 5.8 Mala (22/06/21) and Mw 8.0 Lagunas (26/05/2019) events (Figure 7.b). The maximum strain values are observed in the range of depths between 60 m and 80 m (Figure 7.c), where the first noticeable impedance ratio, from 746 m/s to 929 m/s, is found in the San Miguel profile (Figure 5).

In the case of Villa El Salvador, the RVT-based response spectra fits adequately under the design spectrum and shows good agreement with the results from the time-series analysis [22], though some discrepancy is found for the intermediate period range from 0.50 s and 1.0 s when considering an engineering bedrock type A (Figure 8.a).

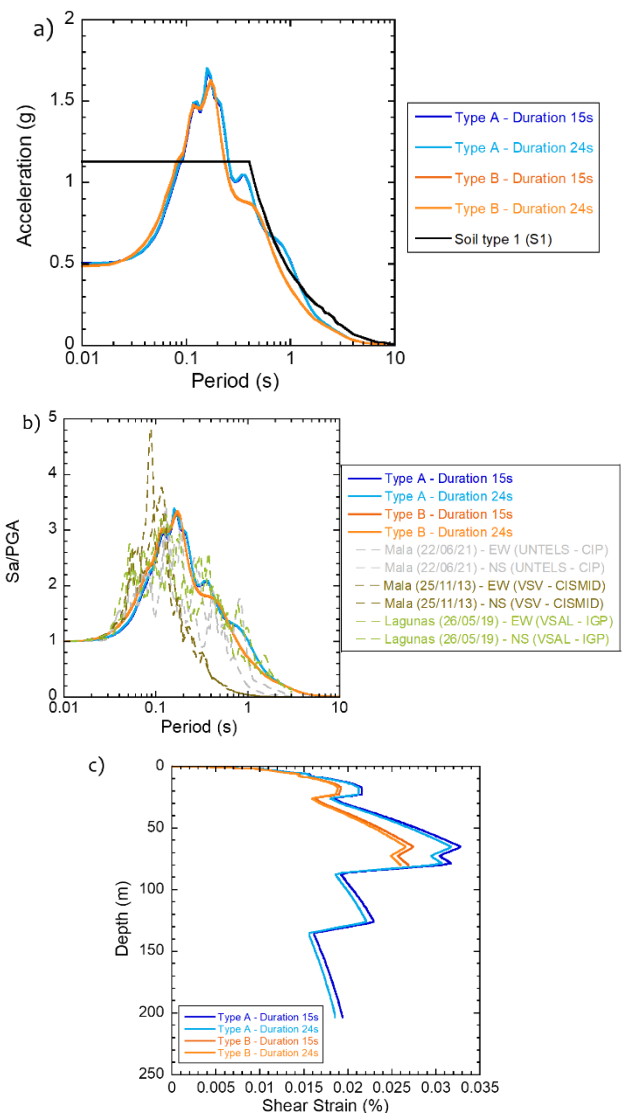


Fig. 7 RVT-based seismic response analysis for San Miguel. a) Response spectra compared with design spectrum. b)

Normalized response spectra compared with those from seismic records. c) Maximum shear strain distribution

The normalized spectrum for the RVT-based analysis shows better agreement with the spectra of the Ml 5.8 Mala (25/11/13) earthquake in the range of periods above 0.15 s, when compared with the results from Mw 8.0 Lagunas (26/06/19) and Ml 5.1 Callao (25/08/21) events. This appears to be reasonable as the hypocentral distance of the event is close to the scenario evaluated from the seismic disaggregation for this study, although the magnitude is significantly smaller. Conversely, spectra for the three events show high amplification values in the low period range which seems to be independent of the epicentral distance and magnitude. Therefore, it is suspected that these apparent high amplifications are due to the normalization of the noise with respect to the low acceleration level recorded in the seismic stations (Figure 8.b). The maximum values of strain are evidenced in the upper 10 m of the profile. Good agreement is observed between the shear strain profile of the RVT-based analysis and the time-series analysis developed in [22] (Figure 8.c). The differences observed between 70 m and 90 m were due the modification of the soil profile to better adjust the Vs profile from Calderon[24].

Finally, La Punta RVT-based analysis performed for the engineering bedrock type B also shows good agreement with the results obtained by Soto[22]. On the other hand, it is observed that amplification factors for the computed spectra exceed those from the design spectra in a period range from 0.60 s to 0.80 s, i.e. around the corner period defined in the Peruvian seismic code (Figure 9.a).

The normalized spectrum for the RVT-based analysis shows better agreement with those for seismic events in the period range between 0.10 s and 0.80 s (Figure 9.b). A sharp peak is observed around 0.25 s for the Ml 5.8 Mala event (22/06/2021) recorded in D855 (CIP) and DSMI (CISMID) stations located about 1.2 km from the study site. This agrees well with the shape of the normalized spectrum from the RVT-based analysis. Even though a high variability in the soil profile of the study area is observed, our response spectra based in RVT can be considered as an envelope of the overall behavior. On the other hand, records for the Mw 8.0 Lagunas earthquake show amplification around 1.40 s, which might be related to the low frequency content caused by this far-field event. On the other hand, all seismic events show high amplification for periods shorter than 0.10 s. This characteristic seems to be independent of the hypocentral distance and magnitude and can be also associated, as in the previous cases, with the normalization of the noise with respect to the low

acceleration level recorded. Finally, the maximum strains are evidenced in the upper 20 m of the profile. Lower strain values were obtained when compared with those from a time-series analysis[22], which might be associated with the implementation of different types of dynamic curves for each study.

As previously observed from the analyses, the influence of the deeper soil substructure is not captured by the design spectra from the Peruvian seismic design standard. Thus, it might be adequate for the soil classification criteria to include additional indexes, such as the fundamental period of vibration. Currently, some efforts have been made on this matter, as can be consulted in Verdugo & Peters[32] and Ruz & Liam[33].

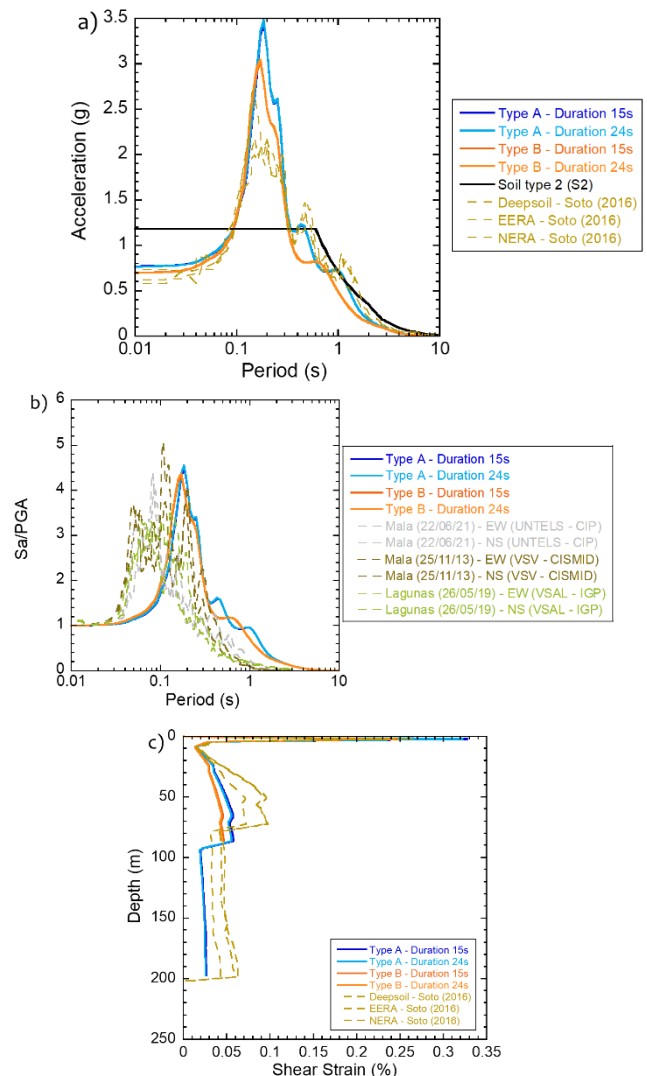


Fig. 8 RVT-based seismic response analysis for Villa El Salvador. a) Response spectra compared with design spectrum and time-history analysis[22]. b) Comparison of normalized response spectra and those from seismic records. c) Maximum shear strain distribution

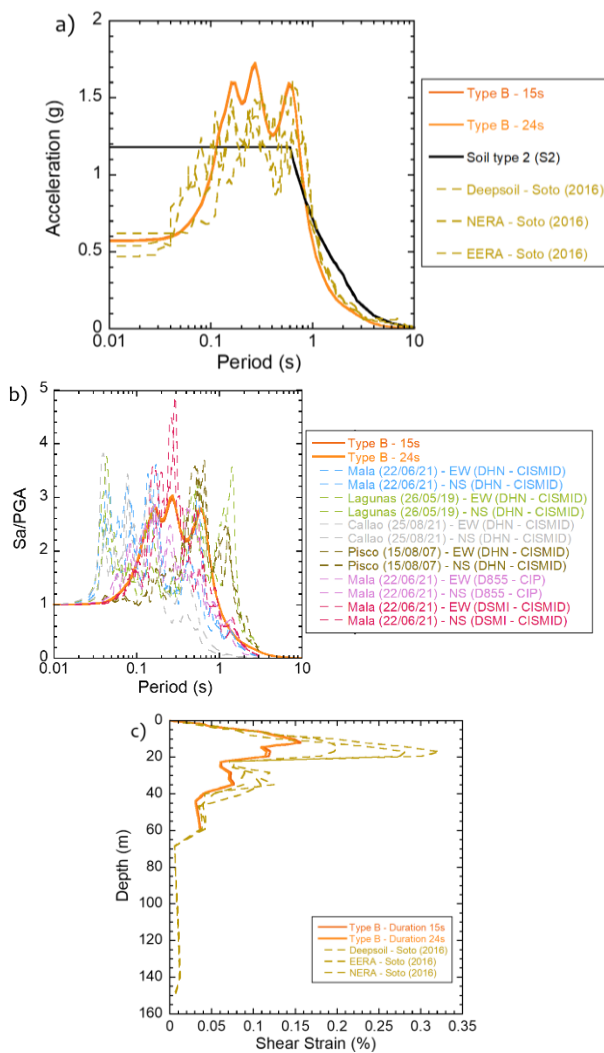


Fig. 9 RVT-based seismic response analysis for La Punta. a) Response spectra compared with design spectrum and time-history analysis[22]. b) Comparison of normalized response spectra and those from seismic records. c) Maximum shear strain distribution

CONCLUSIONS

Random Vibration Theory (RVT)- and time series-based analyses show good agreement in terms of response spectra and maximum shear strain profile shapes. In addition, RVT-based response spectral shapes show consistency with the response spectra from seismic records, despite the differences in the short (below 0.10 s) and long (above 0.80 s) period range that could be related to urban noise and far-field effects, respectively. In general, RVT-based seismic response analysis using the Inverse Random Vibration Theory (IRVT) to define the input motion do not show significant sensitivity of the seismic response spectrum with respect to the duration of the motion for the evaluated sites.

Villa El Salvador and La Punta had available records from nearby stations. In the case of Villa El Salvador, good agreement between results from accelerograms recorded in two different stations was observed, indicating a consistently homogeneous soil deposit. Conversely, seismic records in La Punta displayed different peaks in the acceleration spectra from the stations that recorded the same event, which indicates a variable soil deposit. San Miguel had available records from a station located 3.3 km away from the study site. The evaluated records showed good agreement with the analysis response spectrum, indicating a homogeneous soil deposit. From the obtained results, it is recommended to consider additional parameters to take into consideration the influence of the deeper part of soil profiles in seismic characterization, since indexes, such as v_{s30} , captures the impact of the surficial profile only.

Finally, within the framework of this research, RVT showed to be a viable alternative to perform seismic response analysis as it only requires the implementation of Probabilistic Seismic Hazard Analysis. Results are promising for the evaluated profiles and complementary estimations for profiles with distinct geomorphological characteristics are expected to be evaluated in the future.

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